

# **A new model for polythermal ice**

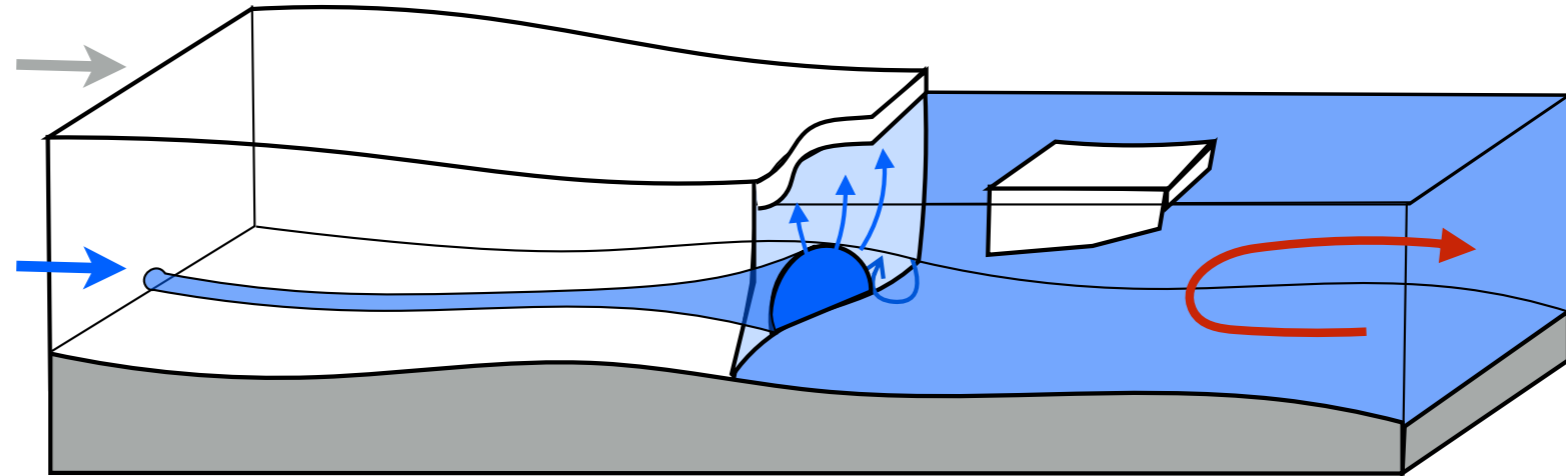
incorporating gravity-driven meltwater drainage

**Ian Hewitt, University of Oxford**

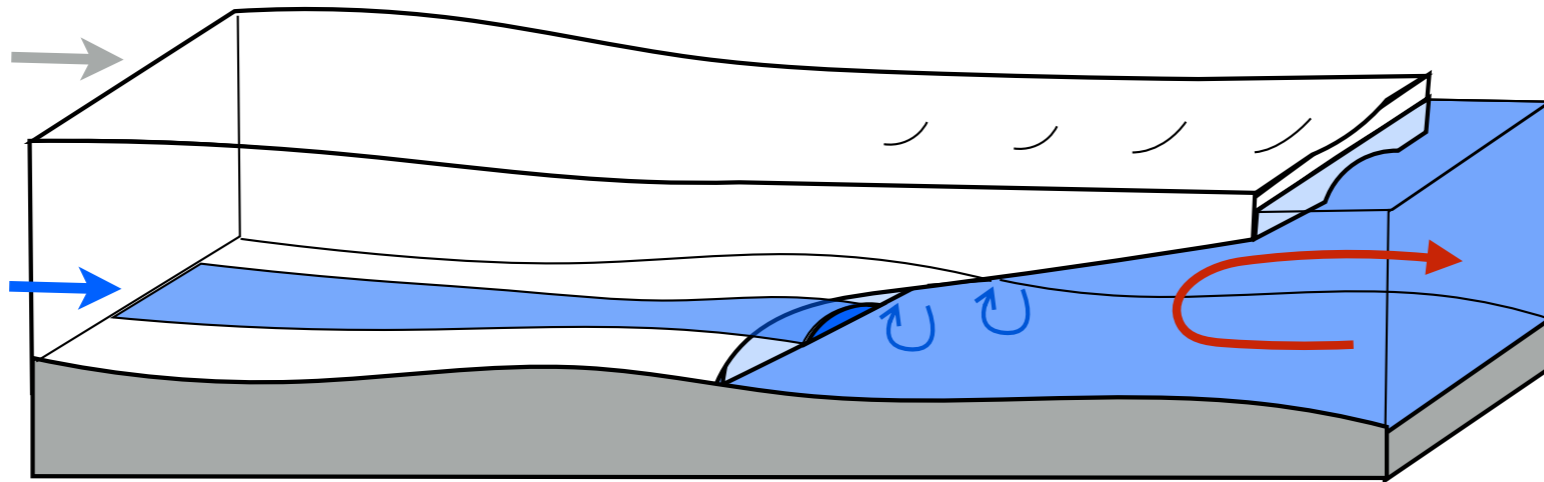
Christian Schoof, University of British Columbia

# From EGU:

I. What do models tell us about how subglacial discharge is delivered at grounding lines?



II. How does the spatial distribution of subglacial discharge affect the shape of ice shelves?



Subglacial channels have a 'trumpet-like' shape near the margin

# **A new model for polythermal ice**

incorporating gravity-driven meltwater drainage

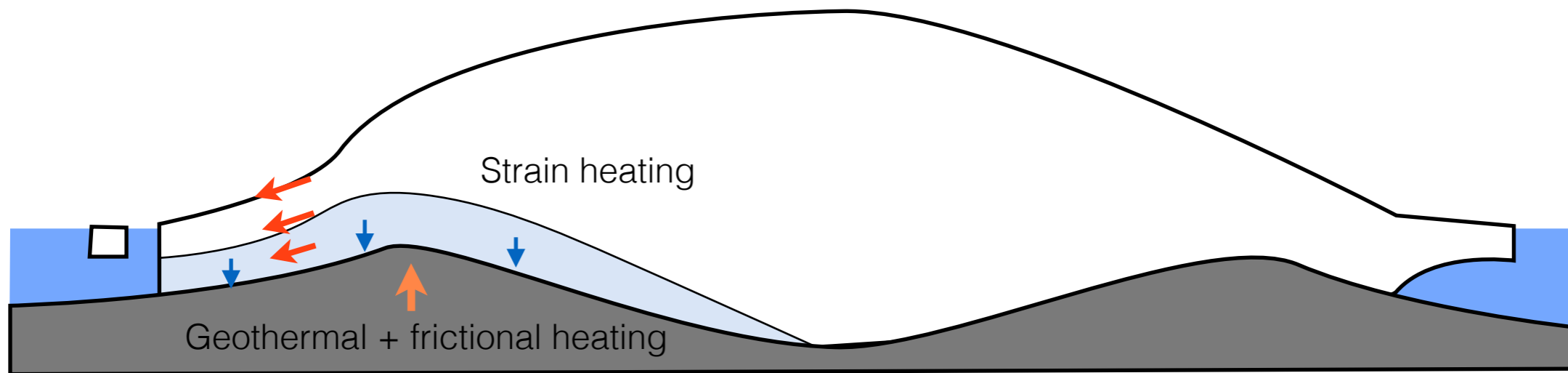
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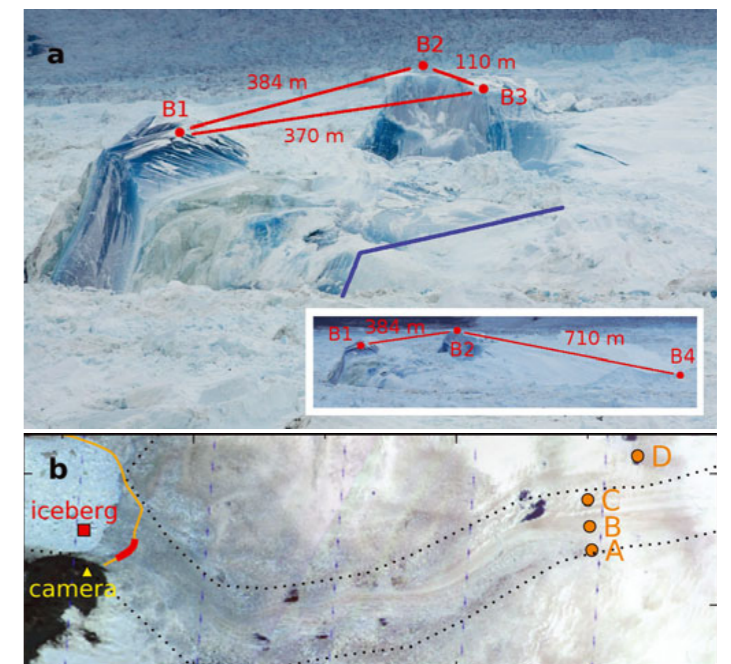
# Motivation

**Goal** - provide a **simple model** that

- predicts temperature and water content of polythermal ice.
- allows water to drain from the ice by porous flow.

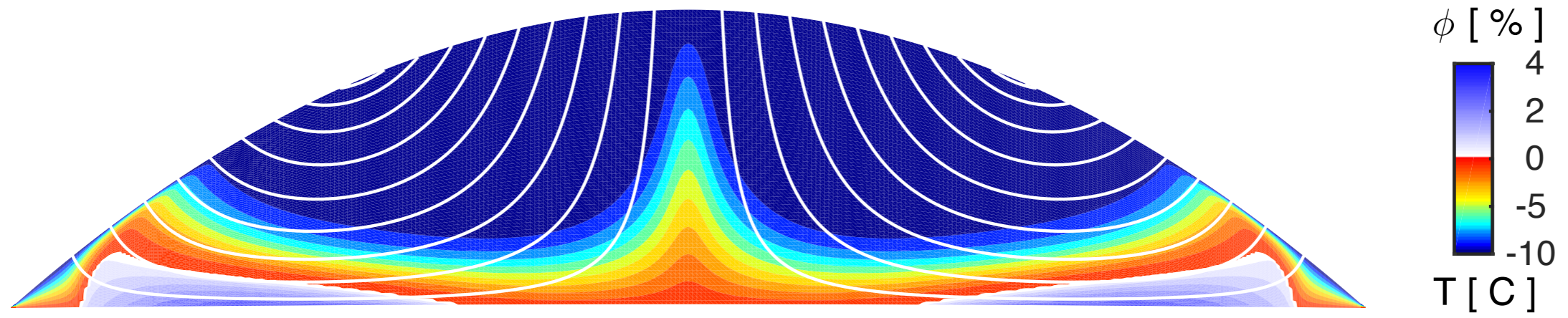


- Ice flow depends on temperature and water content.
- May be fast dynamical feedbacks between water content and ice flow.



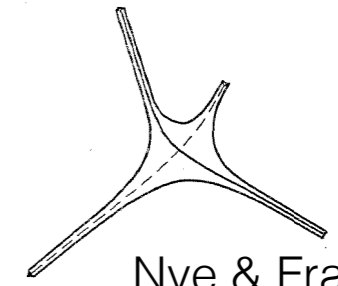
Lüthi et al (2009)

# Motivation

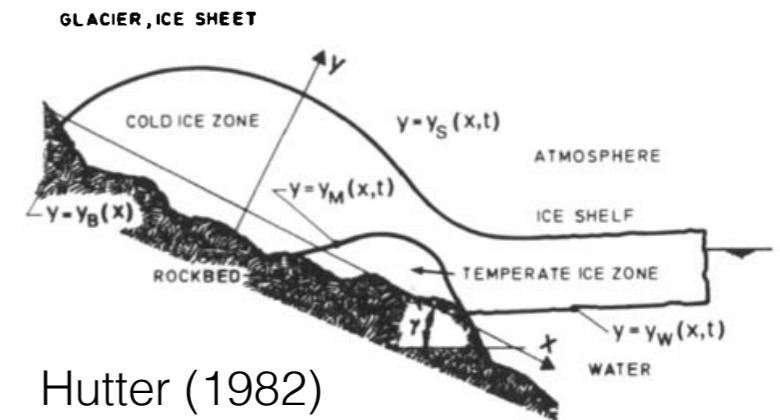


# Previous work on polythermal ice

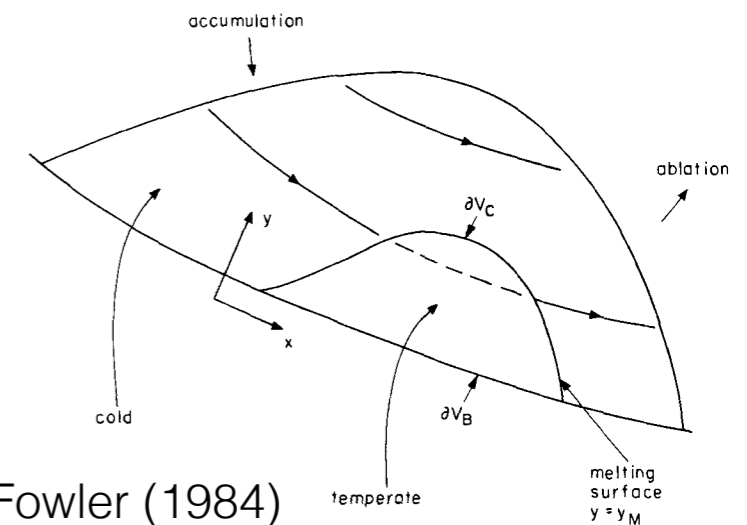
- Theory**
- Lliboutry (1971,1976), Nye & Frank (1973) - permeability
  - Fowler & Larson (1978) - continuum formulation, no moisture movement
  - Hutter (1982) - mixture theory, diffusive moisture transport
  - Fowler (1984) - two-phase theory, Darcy's law for moisture transport



Nye & Frank (1973)



Hutter (1982)

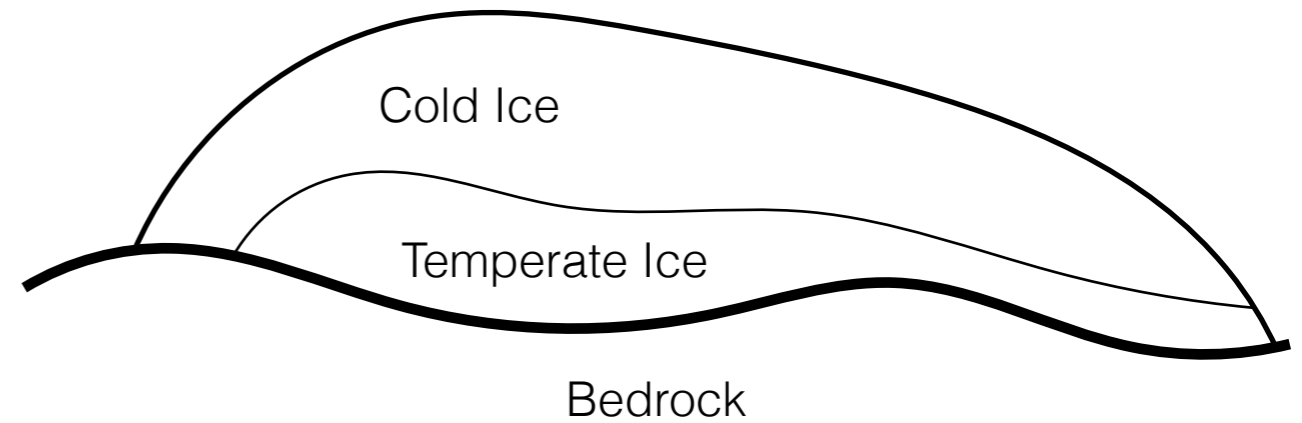


Fowler (1984)

## Computational models

- Greve (1997) - two layer, explicit determination of 'CTS', switch-like drainage function
- Aschwanden et al (2012) - enthalpy gradient method

# Problem formulation



## Stokes flow (or approximation)

$$\nabla \cdot \mathbf{u} = 0$$

$$\frac{\partial \tau_{ij}}{\partial x_j} - \frac{\partial p}{\partial x_i} = -\rho g_i$$

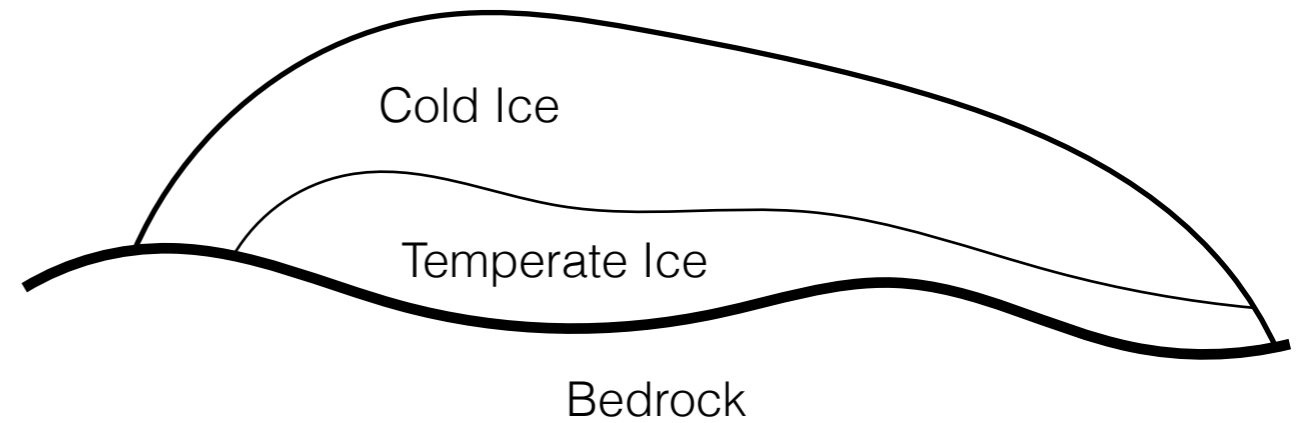
$$\tau_{ij} = A^{-1/n} \dot{\epsilon}^{1/n-1} \dot{\epsilon}_{ij}$$

$$A = A(T, \phi)$$

$\phi$  = water content (porosity)

**This talk** - ice velocity prescribed (decoupled from thermodynamics)

# Problem formulation



## Energy conservation

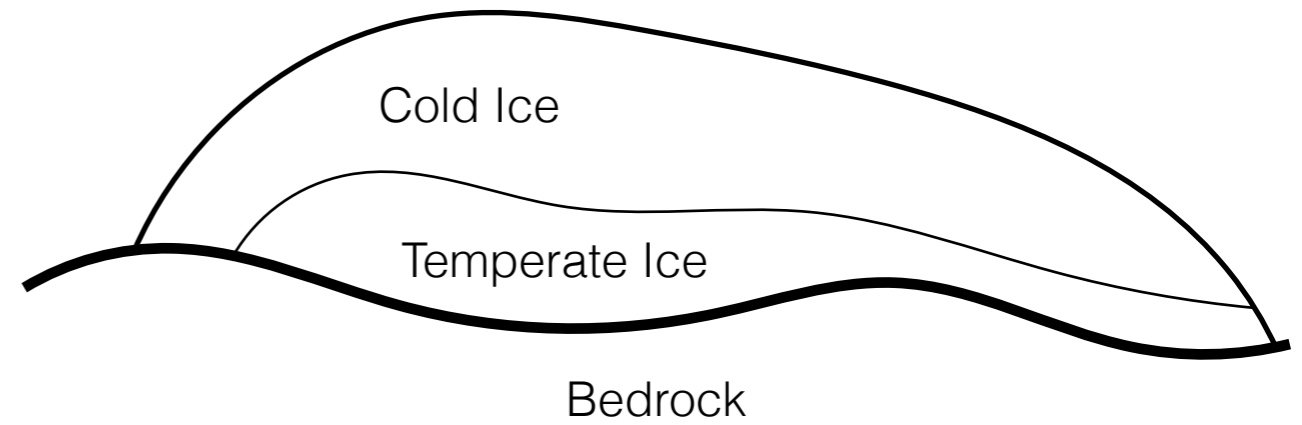
Cold ice  $\rho c \left( \frac{\partial T}{\partial t} + \mathbf{u} \cdot \nabla T \right) = \nabla \cdot (k \nabla T) + \tau_{ij} \dot{\epsilon}_{ij}, \quad \phi = 0, \quad T \leq T_m$

Temperate ice  $\rho_w L \left( \frac{\partial \phi}{\partial t} + \mathbf{u} \cdot \nabla \phi \right) + \rho_w L \nabla \cdot \mathbf{j} = \tau_{ij} \dot{\epsilon}_{ij}, \quad T = T_m, \quad \phi > 0$



Relative water flux  $\mathbf{j} = -\nu \nabla \phi$  - enthalpy gradient method

# Problem formulation



## Energy conservation

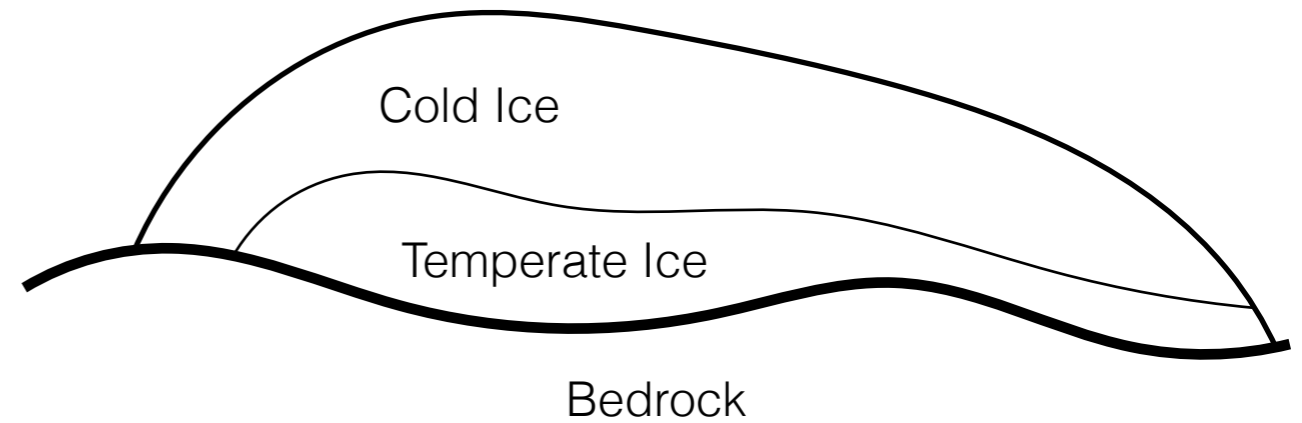
In terms of 'enthalpy',

$$h = \rho c T + \rho_w L \phi$$

$$\left( \frac{\partial h}{\partial t} + \mathbf{u} \cdot \nabla h \right) + \nabla \cdot \mathbf{Q} = \tau_{ij} \dot{\epsilon}_{ij}, \quad \mathbf{Q} = \begin{cases} -k \nabla T & h < \rho c T_m, \\ \rho_w L \mathbf{j} & h \geq \rho c T_m. \end{cases}$$



# Problem formulation



## Energy conservation

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Relative moisture flux  
(Darcy's law)

$$\mathbf{j} = \frac{k_0 \phi^2}{\eta_w} (\rho_w \mathbf{g} - \nabla p_w)$$

Permeability (?)

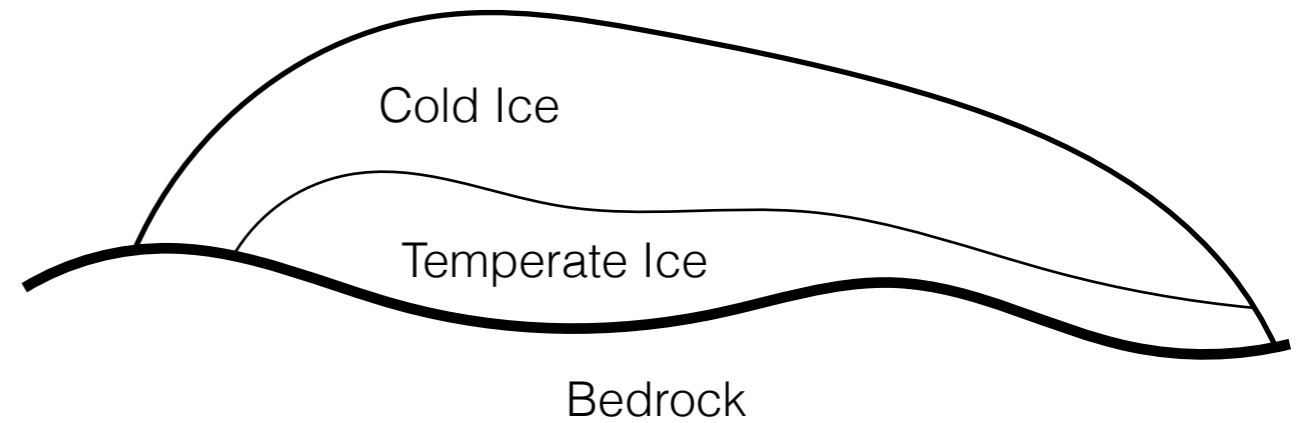
Pore pressure

$$p_w = p - p_e$$

Effective pressure

$$\nabla p \approx \rho \mathbf{g}$$

# Problem formulation



## Energy conservation

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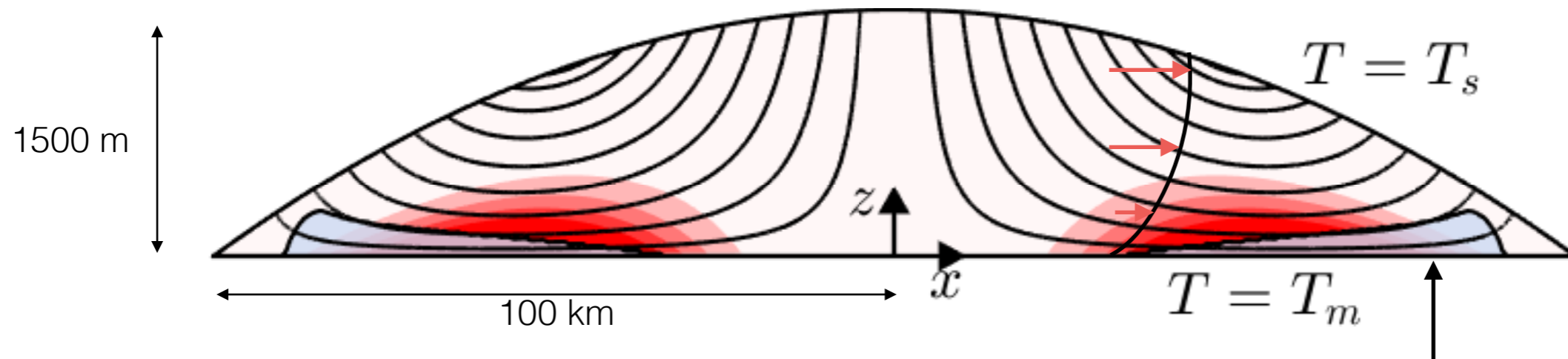
$$\nabla p \approx \rho \mathbf{g}$$

Viscous compaction  
e.g. Hewitt & Fowler 2008

$$\nabla \cdot \mathbf{j} = \frac{\phi p_e}{\eta}$$

$$\mathbf{j} = \frac{k_0 \phi^\alpha}{\eta_w} ((\rho_w - \rho) \mathbf{g} + \nabla p_e)$$

# Ice divide example

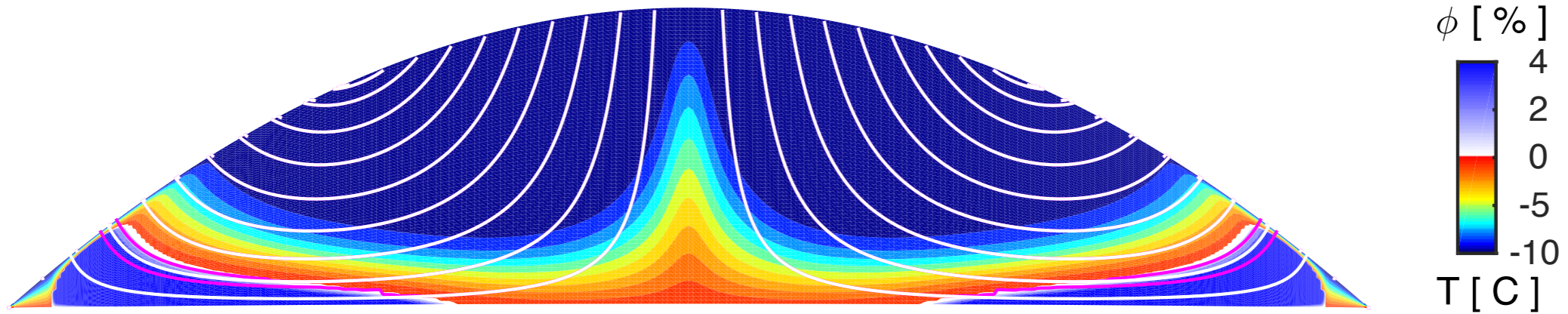


Pore pressure = subglacial drainage pressure  $p_e = N_0$

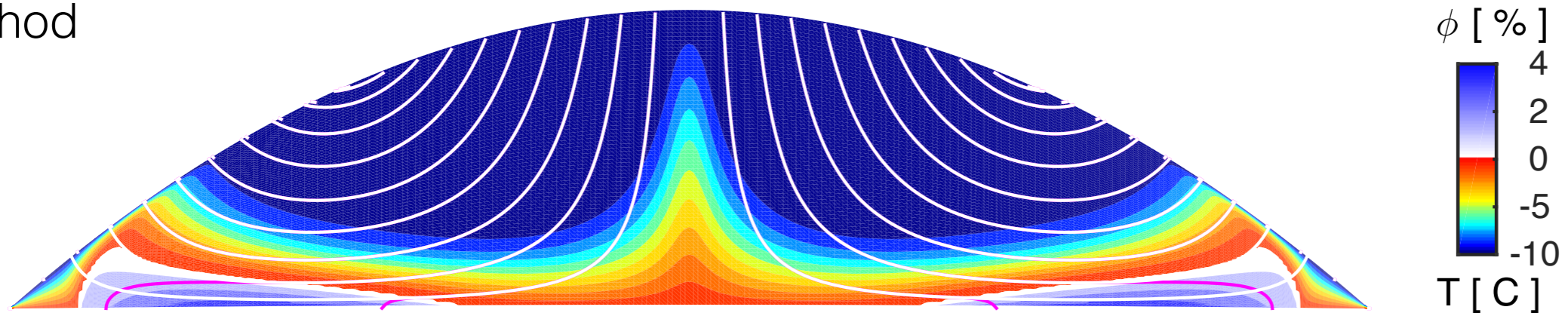
- Velocity from shallow ice approximation (thermodynamically decoupled).

# Ice divide example

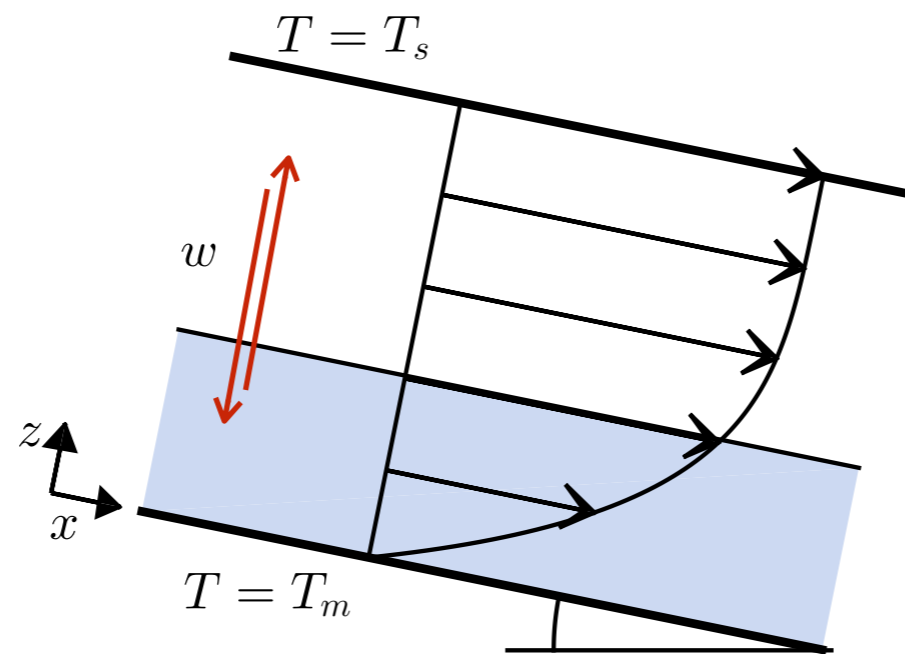
'Standard' enthalpy  
gradient method



Compaction pressure  
method



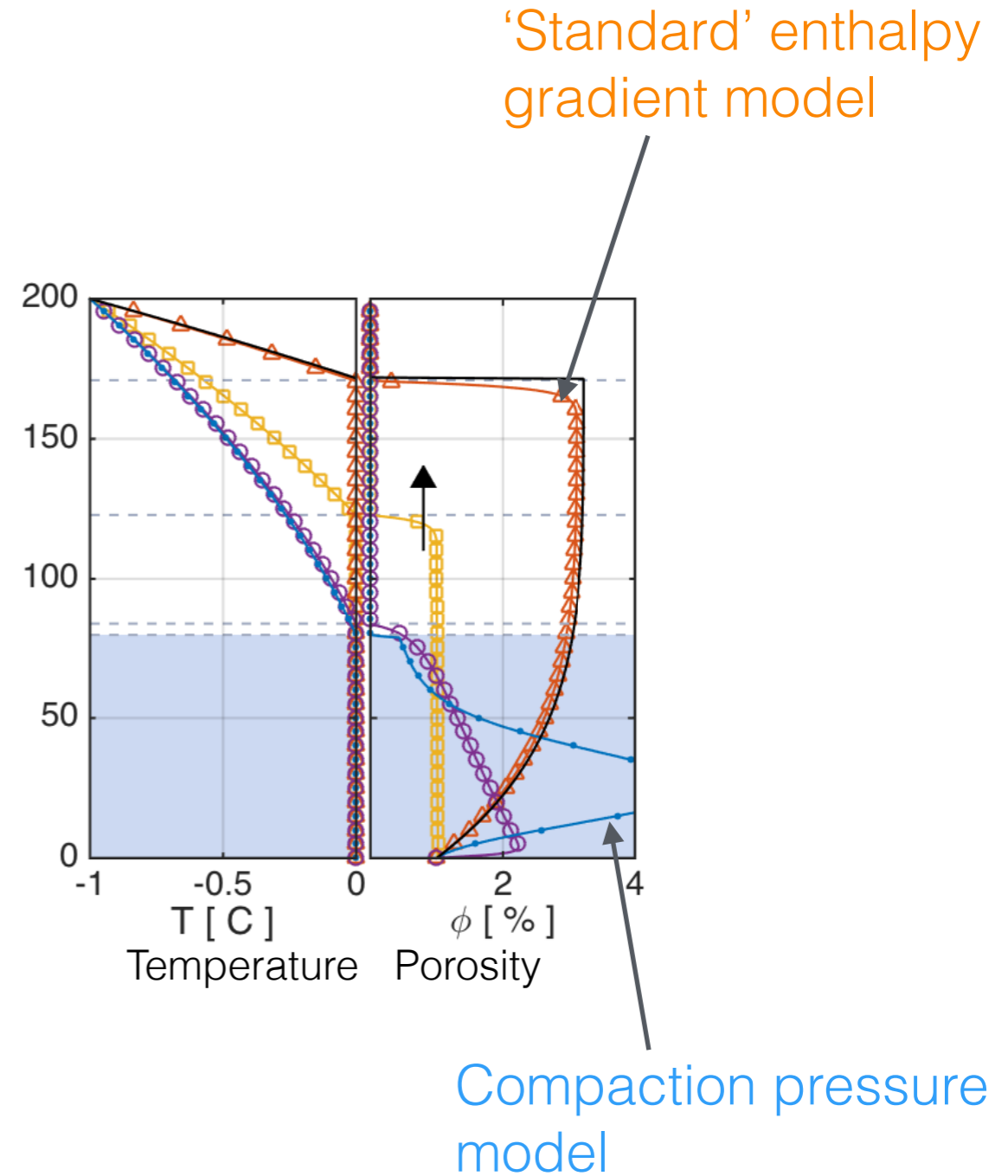
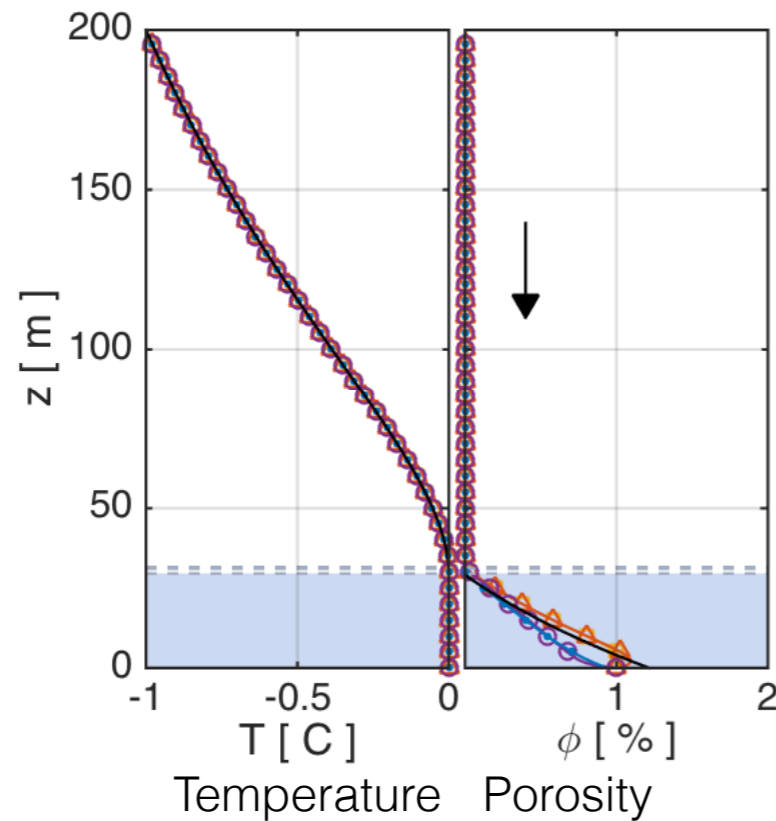
# Slab glacier test case



- Greve & Blatter (2009), Kleiner et al (2015), Blatter & Greve (2015)

# Slab glacier test case

- Comparison of different models

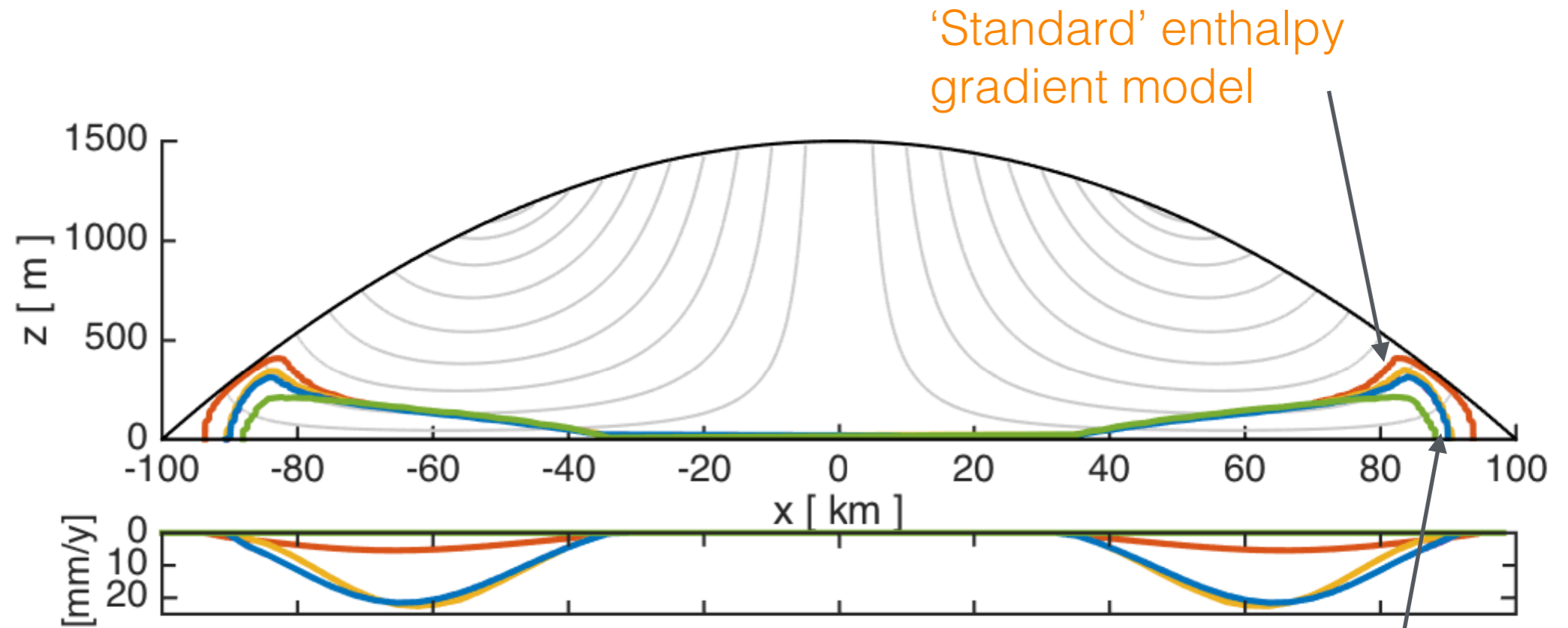


- Standard enthalpy gradient well approximates no-water-transport solution.
- Gravity-driven drainage results in less temperate ice - and some interesting behaviour...

# Ice divide example

Cold-temperate transition

Meltwater flux to bed



'Standard' enthalpy gradient model

Compaction pressure model

# Summary

- Suggested a **simple model** to incorporate polythermal ice in existing ice-sheet models - alternative to enthalpy gradient method.
- Model allows **water transport** through the ice, and connection with subglacial drainage - but more knowledge of permeability needed.
- Worth exploring dynamics of temperate ice + subglacial water + sediments further.